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Influence of Leaf Area Index on the Radiometric Resistance to Heat Transfer

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Abstract Sensible heat flux is an important component of the surface energy balance. Land surface models often use the radiative surface temperature instead of the aerodynamic temperature to predict the surface sensible flux, because the former is much easier to observe by remote sensing or to compute from the surface energy balance equation. Here, measurements from 44 FLUXNET sites are used to examine the stability and leaf area index (*LAI*) dependence of the radiometric resistance, a resistance that should be included in the bulk transfer method if the radiometric temperature is used for the flux calculation. Results show that the radiometric resistance is much higher under stable conditions than under unstable conditions. In unstable conditions, the radiometric resistance is highly sensitive to *LAI*, decreasing exponentially as *LAI* increases. Omission of the radiometric resistance from the bulk transfer method causes a large overestimation in the sensible heat flux, especially for low-*LAI* surfaces and under unstable conditions.

Keywords Aerodynamic temperature · Leaf area index · Radiative surface temperature · Radiometric resistance · Sensible heat flux

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1 Introduction

The bulk transfer equation (Garratt and Hicks 1973; Monteith 1973; Garratt and Francey 1978) is a widely employed method for estimating sensible heat flux (H) in land-surface models (LSMs) (Mahrt and Vickers 2004), i.e.,

$$H = \rho C_p \frac{\theta_{0h} - \theta_a}{R_{ah}}, \tag{1}$$

where ρ is air density, C_p is the specific heat capacity of air at constant pressure, θ_a is the potential temperature at a reference height in the surface layer, θ_{0h} is the effective aerodynamic temperature of the surface, and R_{ah} is a heat transfer resistance for the pathway from θ_{0h} to θ_a. The heat transfer resistance, R_{ah}, can be formulated as

$$R_{ah} = \frac{\left[\ln\left(\frac{z-d}{z_0}\right) - \Psi_M\right]\left[\ln\left(\frac{z-d}{z_h}\right) - \Psi_H\right]}{k^2 u}, \tag{2}$$

where k is the von Karman constant, z is the measurement height of θ_a and wind speed u, d is the displacement height, z_0 is the aerodynamic roughness length, z_h is the thermal roughness length, and Ψ_M and Ψ_H are the stability correction functions for momentum and heat, respectively. Note that z_0 and z_h are profile characteristics. The implicit assumption is that the standard Monin–Obukhov functions are valid down to the top of the roughness sublayer, but are invalid below this level.

The accuracy of Eq. 1 depends on how the surface temperature is chosen. Use of R_{ah} defined in Eq. 2 requires θ_{0h} to be the effective aerodynamic temperature at $z = d + z_h$. The two roughness lengths in Eq. 2 are typically different (Garratt and Hicks 1973), and their ratio is a function of surface stiffness and the roughness Reynolds number (Molder and Lindroth 2001). A universal and robust parametrization of z_h across a large variety of land-cover types does not exist, and so two alternative approaches are used to remove z_h from the bulk transfer method. In one approach, z_h is assumed to be equal to z_0, and R_{ah} is simply reduced to the aerodynamic resistance R_a,

$$R_a = \frac{\left[\ln\left(\frac{z-d}{z_0}\right) - \Psi_M\right]\left[\ln\left(\frac{z-d}{z_0}\right) - \Psi_H\right]}{k^2 u}, \tag{3}$$

while in the second approach, R_{ah} is expressed as a combination of the aerodynamic resistance, R_a, and the excess resistance, R_{ex},

$$R_{ah} = R_a + R_{ex}. \tag{4}$$

In neutral stability, manipulation of Eqs. 2–4 gives

$$R_{ex} = \frac{\ln(z_0/z_h)}{k u_*}, \tag{5}$$

where u_* is the friction velocity (Garratt and Hicks 1973).

One difficulty with the above formulations is that θ_{0h} cannot be measured directly. In field campaigns, this temperature is usually inferred from extrapolation of the temperature profile to z_0 or z_h according to Monin–Obukhov similarity theory (Thom et al. 1975; Garratt and Francey 1978). Although z_0 and z_h are profile characteristics, the corresponding roughness elements typically lie in the roughness sublayer, where the similarity theory does not hold. Therefore, this extrapolation of the temperature profile is an important source of

error. In addition, vertical temperature profiles are generally not available in remote-sensing applications (Kustas et al. 2003; Matsushima 2005; Kustas and Anderson 2009; Mu et al. 2011). For these reasons, in practice, θ_{0h} is usually approximated by the surface radiative temperature θ_s, which can be measured by remote-sensing techniques or computed from the surface energy balance equation (Sun and Mahrt 1995). Use of θ_s in place of θ_{0h} for the bulk parametrization is especially attractive to modellers, because in almost all land-surface models, θ_s is an important prognostic variable derived from solution of the surface energy balance equation.

Previous researchers have reported that θ_s and θ_{0h} can be very different, in both stable and unstable conditions, and a simple substitution of θ_s for θ_{0h} can result in systematic high biases in the prediction of sensible heat flux (Choudhury et al. 1986; Beljaars and Holtslag 1991; Hall et al. 1992). To date, efforts to remediate this problem have fallen into three categories:

- Prediction of θ_{0h} from θ_s. In some studies, the relationship between the radiative and the effective aerodynamic temperature is expressed as a function of z_h (Garratt and Francey 1978; Brutsaert and Sugita 1992). Beljaars and Holtslag (1991) suggested that the relationship should depend on the temperature scale θ_* (defined as $H/\rho C_p u_*$) and thermal stability. Mahrt and Vickers (2004) and Matsushima (2005) found that $\theta_s - \theta_{0h}$ is related to solar radiation and leaf area index (*LAI*).

- Introduction of a new roughness length z_r. Here, z_n is called the "radiometric roughness length," being smaller than z_h by up to a few orders of magnitude. In a number of studies, the difference between the radiometric roughness length z_r and z_0 is represented using a redefined kB^{-1} parameter $\left(kB^{-1} = \ln \frac{z_0}{z_r}\right)$, instead of the standard formulation ($kB^{-1} = \ln \frac{z_0}{z_h}$); this modified kB^{-1} parameter is related to the roughness Reynolds number and θ_s (Kustas et al. 1989; Sugita and Brutsaert 1990; Stewart et al. 1994; Sun and Mahrt 1995; Troufleau et al. 1997; Verhoef et al. 1997; Yang et al. 2003) or is physically modelled based on the "localized near-field" Lagrangian theory (McNaughton and Vandenhurk 1995; Su et al. 2001). These studies show that kB^{-1} behaves erratically, and that a universal parametrization may be unattainable.

- To add an extra resistance, R_r: this "radiometric resistance" is added to the bulk transfer formulation (Lhomme et al. 1988; Stewart et al. 1994), such that

$$H = \rho C_p \frac{\theta_s - \theta_a}{R_r + R_{ah}}, \tag{6}$$

where $R_{ah} = R_a + R_{ex}$. According to Lhomme et al. (1988) and Stewart et al. (1994), R_r is generally larger than R_{ah}.

A number of studies have suggested the importance of *LAI* and thermal stability in controlling the biases caused by substituting θ_s for θ_{0h}. Beljaars and Holtslag (1991) found that $\theta_{0h} - \theta_s$ varies from 6 K in stable conditions to -6 K in unstable conditions. Both Troufleau et al. (1997) and Verhoef et al. (1997) found that, for sparse vegetation (low *LAI*), kB^{-1} shows a large mean value and a large variation. Mahrt and Vickers (2004) reported a significant positive linear dependence of $\theta_s - \theta_{0h}$ on solar radiation and a negative dependence on *LAI*. Solar radiation partly influences the heat transfer through its effect on thermal stability. These relationships have been established with observations over several surface types, including a bare-soil site, three grassland sites, a cropland site, and four forest sites. Kustas et al. (2007) found that the slopes and intercepts of the linear regressions that Mahrt and Vickers (2004) proposed vary considerably when applied to a wider range of landscapes (*LAI* 0.5–3.0) and meteorological conditions. Similarly, the correction factor proposed by Matsushima (2005) for $\theta_s - \theta_{0h}$ has a strong correlation with *LAI*.

Recently, Zheng et al. (2012) developed new formulations of the roughness lengths using a green vegetation fraction derived from the remotely sensed normalized difference vegetation index, the latter being known to be a good estimator of *LAI*. The stability effects on the radiometric resistance R_r or kB^{-1} have been somewhat overlooked in previous studies (Stewart et al. 1994; Troufleau et al. 1997; Matsushima 2005); for example, Matsushima (2005) showed that the effects of atmospheric stability in the surface layer can be ignored in the parametrized radiometric roughness length or temperature "correction" factors.

The primary objective of the present study is to revisit the roles of *LAI* and thermal stability in the bulk formulation (Eq. 6). Our focus is on patterns of the radiometric resistance R_r across a diverse set of ecosystems and local climates. Previous studies have been restricted to a small number of sites (one to nine). Here, we deploy data obtained from 44 sites in the FLUXNET network. Previous reports have confirmed a large variation in hourly kB^{-1} values (Troufleau et al. 1997; Verhoef et al. 1997), partly due to the high sensitivity of kB^{-1} to measurement errors in the micrometeorological variables (Verhoef et al. 1997). Here, we examine how mean resistance varies across the sites, instead of using 30-min values at a specific site as done in previous studies. The advantage of this site-mean approach is that it identifies patterns caused by land-cover differences, compared with synoptic meteorological fluctuations and instrument noise. In addition, we provide an assessment of the prediction error in sensible heat flux caused by omission of R_r from Eq. 6.

2 Data and Methods

2.1 Surface Observations

We examine tower observations from 44 FLUXNET sites in the USA, Canada, and China (Table 1). FLUXNET is a global network of surface eddy-covariance towers, with the observations covering a large range of u_* (site annual mean value 0.1–1.5 m s^{-1}). Although u_* is related to wind speed, the annual mean value can be viewed as a proxy of the surface roughness, having a high correlation with z_0 (linear correlation = 0.60). We chose these sites because measurements of the longwave components of the surface radiation balance were available, in addition to the energy and momentum fluxes and micrometeorological variables, and for a sufficiently long time (>12 months). They span a large range of *LAI* (0 to 9) and canopy height (0.1 to 33 m), and are divided into four categories: 19 conifer forest sites (with nearly constant *LAI* throughout the year), 12 deciduous forest sites (with seasonal *LAI* change), 6 grassland sites (with seasonal *LAI* change), and 7 cropland sites (with large *LAI* changes between growing and nongrowing seasons). For each site, we selected one year of measurements and confined our analysis periods to December to February, and June to August, for the winter and summer seasons (>90 % data coverage for each site for each season). We also separate the data into stable and unstable conditions. The original data are 30-min values for 40 sites; for four sites (site IDs: US-MMS, US-Ne1, US-Ne2, and US-Ne3; Table 1) only hourly measurements are available.

Except for the Mead irrigated cropland site in Nebraska (site ID US-Ne1, Table 1), we report all results as site mean values. At the Mead cropland site, high-frequency measurements of *LAI* were available from 2006 to 2012, showing large *LAI* changes within the growing season. To capture these seasonal variations, we analyzed the daily mean quantities for this site.

Table 1 Selected FLUXNET sites and their land-surface types

Site ID	Site name	State/prov.	Lat. (N)	Long.	Vegetation	Reference
CA-WP1	Western peatland	Alberta	54.95	−112.47	Conifer	Syed et al. (2006)
CA-Obs	Old black spruce	Saskatchewan	53.99	−105.12	Conifer	Jarvis et al. (1997)
CA-SJ2	SK-2002 Jack pine	Saskatchewan	53.94	−104.65	Conifer	Coursolle et al. (2006)
CA-Ojp	Old Jack pine	Saskatchewan	53.92	−104.69	Conifer	Baldocchi et al. (1997)
CA-SJ3	SK-1975 (Young) Jack pine	Saskatchewan	53.88	−104.65	Conifer	Mahrt and Vickers (2002)
CA-Oas	Old aspen	Saskatchewan	53.63	−106.20	Deciduous	Blanken et al. (1997)
CA-Ca1	Douglas fir 1949	British Columbia	49.87	−125.33	Conifer	Humphreys et al. (2003)
CA-Qfo	Quebec mature boreal forest	Québec	49.69	−74.34	Conifer	Bergeron et al. (2007)
CA-Ca3	Douglas fir 1988	British Columbia	49.53	−124.90	Conifer	Humphreys et al. (2006)
US-FPe	Fort Peck	Montana	48.31	−105.10	Grassland	Schmidt et al. (2011)
CA-Gro	Groundhog River mixedwood	Ontario	48.22	−82.16	Mixed	Coursolle et al. (2006)
CA-Na1	Nashwaak Lake 1 1967 balsam fir	New Brunswick	46.47	−67.10	Deciduous	Coursolle et al. (2006)
US-WCr	Willow Creek	Wisconsin	45.81	−90.08	Deciduous	Davis et al. (2003)
US-UMB	UMBS	Michigan	45.56	−84.71	Deciduous	Schmid et al. (2003)
CA-Mer	Mer Bleue eastern peatland	Ontario	45.41	−75.52	Deciduous	Coursolle et al. (2006)
US-Ho2	Howland Forest West Tower	Maine	45.21	−68.75	Conifer	Thornton et al. (2002)
US-Ho3	Howland Forest East Tower	Maine	45.21	−68.73	Conifer	Hollinger et al. (2004)
US-Ho1	Howland Forest Main	Maine	45.20	−68.74	Conifer	Hollinger et al. (1999)
US-MRf	Marys River	Oregon	44.65	−123.55	Conifer	Vickers et al. (2009)
US-Me2	Metolius intermediate pine	Oregon	44.45	−121.56	Conifer	Law et al. (2004)
US-Bkg	Brookings	South Dakota	44.35	−96.84	Grass/crops	Gilmanov et al. (2005)
US-Me3	Metolius second young pine	Oregon	44.32	−121.61	Conifer	Vickers et al. (2009)
CN-Du2	Duolun-grassland	Inner Mongolia	42.05	116.28	Grassland	Chen et al. (2009)
US-GLE	GLEES	Wyoming	41.36	−106.24	Conifer	Massman (2000)

Table 1 continued

Site ID	Site name	State/prov.	Lat. (N)	Long.	Vegetation	Reference
US-Ne3	Mead rainfed maize–soybean rotation	Nebraska	41.18	−96.44	Cropland	Suyker et al. (2005)
US-Ne1	Mead irrigated continuous maize	Nebraska	41.17	−96.48	Cropland	Suyker et al. (2004)
US-Ne2	Mead irrigated maize–soybean rotation	Nebraska	41.16	−96.47	Cropland	Suyker et al. (2004)
US-NR1	Niwot Ridge	Colorado	40.03	−105.55	Conifer	Monson et al. (2002)
US-Bo1	Bondville	Illinois	40.01	−88.29	Cropland	Hollinger et al. (2005)
US-Slt	Silas Little Experimental Forest	New Jersey	39.91	−74.60	Deciduous	Clark et al. (2010)
US-MMS	Morgan Monroe State Forest	Indiana	39.32	−86.41	Deciduous	Schmid et al. (2000)
US-CaV	Canaan Valley	West Virginia	39.06	−79.42	Cropland	Hollinger et al. (2010)
US-MOz	Missouri Ozark	Missouri	38.74	−92.20	Deciduous	Gu et al. (2006)
US-Var	Vaira Ranch	California	38.41	−120.95	Shrubs/grassland	Ma et al. (2007)
US-ARM	ARM Southern Great Plains	Oklahoma	36.61	−97.49	Grass/Crops	Fischer et al. (2007)
US-WBW	Walker Branch	Tennessee	35.96	−84.29	Deciduous	Wilson and Meyers (2001)
US-NC1	North Carolina Clearcut	North Carolina	35.81	−76.71	Deciduous	Noormets et al. (2010)
US-NC2	North Carolina Loblolly Plantation	North Carolina	35.80	−76.67	Conifer	Noormets et al. (2010)
US-Fwf	Flagstaff Wildfire	Arizona	35.45	−111.77	Shrubs	Dore et al. (2008)
US-Fmf	Flagstaff Managed Forest	Arizona	35.14	−111.73	Conifer	Dore et al. (2008)
US-Fuf	Flagstaff Unmanaged Forest	Arizona	35.09	−111.76	Conifer	Dore et al. (2008)
US-Goo	Goodwin Creek	Mississippi	34.25	−89.87	Deciduous	Wilson and Meyers (2007)
US-SRM	Santa Rita Mesquite	Arizona	31.82	−110.87	Shrubs/grassland	Scott et al. (2009)
US-Aud	Audubon Research Ranch	Arizona	31.59	−110.51	Grassland	Krishnan et al. (2012)

Fig. 1 Schematic of the resistance decomposition model for calculating the sensible heat flux, H. θ_{0h} is the aerodynamic temperature at the thermal roughness height; θ_0 is the aerodynamic temperature at the momentum roughness height
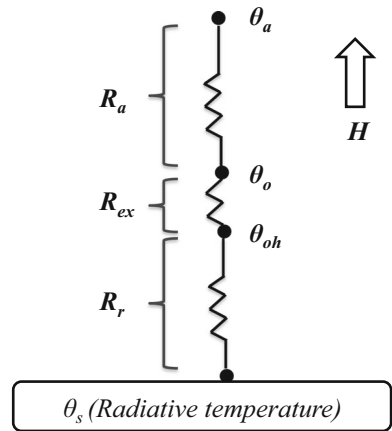
2.2 Data Analysis

We computed the half-hourly or hourly R_r using Eq. 6, where the total heat resistance consists of three additive components (Fig. 1). The aerodynamic resistance R_a was determined with Eq. 3, and the excess resistance R_{ex} was approximated by

$$R_{ex} = \frac{2}{ku_*},\tag{7}$$

which indicates that $z_0/z_h = 7.4$, a typical ratio for rough surfaces (Garratt 1994). The corresponding kB^{-1} value is 2. Making use of Eqs. 4, 6, and 7, we obtain an expression for computing R_r,

$$R_r = \frac{\rho C_p\,(\theta_s - \theta_a)}{H} - R_a - \frac{2}{ku_*}.\tag{8}$$

In Eqs. 2, 3, and 8, H, θ_a, and u_* were available from direct tower measurements, while θ_s, ρ, R_a, and the stability correction factors Ψ_M and Ψ_H were calculated using measured variables, and d and z_0 were assumed as 70 % and 10 % of the canopy height (h), respectively. Note that Ψ_M and Ψ_H were calculated according to the Businger–Dyer formulation, though we tested other formulations of the stability correction functions (Brutsaert 1992), finding that the results obtained were insensitive to the choice of stability correction functions.

We separated the half-hourly or hourly data into stable and unstable conditions using the sign of the surface sensible heat flux. All observations with $H < -5\,\mathrm{W\,m^{-2}}$ were grouped into stable conditions, and all observations with $H > 5\,\mathrm{W\,m^{-2}}$ were grouped into unstable conditions. Observations with $-5 < H < 5\,\mathrm{W\,m^{-2}}$ were considered as neutral or near-neutral conditions and were not used.

The surface radiative temperature T_s was determined from the upward longwave radiation flux and corrected for the surface reflection of the downward longwave radiation flux, as

$$T_s = \left[\frac{L_\uparrow - (1 - \varepsilon)\,L_\downarrow}{\varepsilon\sigma}\right]^{1/4},\tag{9}$$

where $L_\uparrow$ is the upward longwave radiative flux, $L_\downarrow$ is the downward longwave radiative flux, ε is the surface emissivity (assumed to be 0.98), and σ is the Stefan–Boltzmann constant.

3 Results and Discussion

3.1 Relations of Heat Resistances to *LAI* Under Unstable Conditions

Here, we provide a detailed examination of the resistances under unstable conditions. Previous studies have shown that the bulk formulation is much more uncertain under unstable conditions than under stable conditions. The excess resistance R_{ex} shows no obvious dependence on *LAI* (Fig. 2e, f). This resistance arises from the fact that heat is transferred by molecular diffusion through the laminar boundary layer on foliage in immediate contact with the surface, whereas momentum exchange is more efficient due to the existence of both viscous shear and form drag (Thom 1972). The insensitivity of R_{ex} to *LAI* suggests that R_{ex} is controlled primarily by these processes at the leaf scale and is not affected by the amount or arrangement of foliage elements at the canopy scale.

The radiometric resistance R_r shows a much stronger relationship with *LAI* than does R_{ex}, with an exponential decay as *LAI* increases (Fig. 2a, b). For small *LAI* values ($LAI < 1$), R_r is 3.7 times larger than R_{ex} in the summer and 1.4 times larger in the winter. An alternative interpretation is that the effective z_r is much smaller than z_h for sparse vegetation. Because at sites of larger R_r the temperature difference $\theta_s - \theta_{0h}$ should be larger in magnitude, our results are consistent with those of Mahrt and Vickers (2004). Using airborne measurements from the Southern Great Plains Experiment, these authors showed that $\theta_s - \theta_{0h}$ can be described by the simple linear model

$$\theta_s - \theta_{0h} = C\left[S - C_s\left(LAI - LAI_{\text{ref}}\right)\right], \tag{10}$$

where C and C_s are positive regression coefficients, S is the solar radiative flux, and the reference *LAI* value (LAI_{ref}) is approximately 1. This expression illustrates that $\theta_s - \theta_0$ is large when *LAI* is small. They also acknowledged that a more complex function of *LAI* is required for very large *LAI*. The dependence on S is explained by the fact that sunlit surfaces warm up more rapidly than the air at the roughness height.

The aerodynamic resistance R_a behaves similarly to R_r as *LAI* changes (Fig. 2c, d), and in terms of magnitude, R_a is comparable to R_r. The physical explanation of the exponential correlation between the *LAI* and the aerodynamic resistance is connected to convection efficiency. High *LAI* usually corresponds to aerodynamically rough surfaces; these surface can trigger large coherent eddies that are efficient in heat convection, whereas low-*LAI* surfaces (grassland or fallow cropland) are more like bluff-rough surfaces that are less efficient in generating energetic eddies (Stewart et al. 1994; Voogt and Grimmond 2000).

3.2 Summer Versus Winter

There are seasonal differences in the three resistances (Table 2), with Fig. 2 showing that the general relationships with *LAI* are similar between the summer and winter seasons. In the winter, when *LAI* is smaller, the site-mean values of R_r and R_a are more scattered than in the summer (Fig. 2a–d). One reason is related to deciduous forest sites, which have low *LAI* in the winter, but are still aerodynamically rough. Instead of assuming that z_0 is proportional to stand height, an improved parametrization (e.g., Raupach 1994) where z_0 is a function of both canopy height and *LAI*, may reduce the scatter.

In terms of magnitude, R_r shows higher values in the winter than in the summer, and the other two resistances do not show strong seasonal differences. The R_r seasonality is partly explained by *LAI* changes. For the deciduous forest group, the average winter *LAI* is low (1.5), and the average R_r is 66 s m^{-1}. In the summer, when the average *LAI* increases to
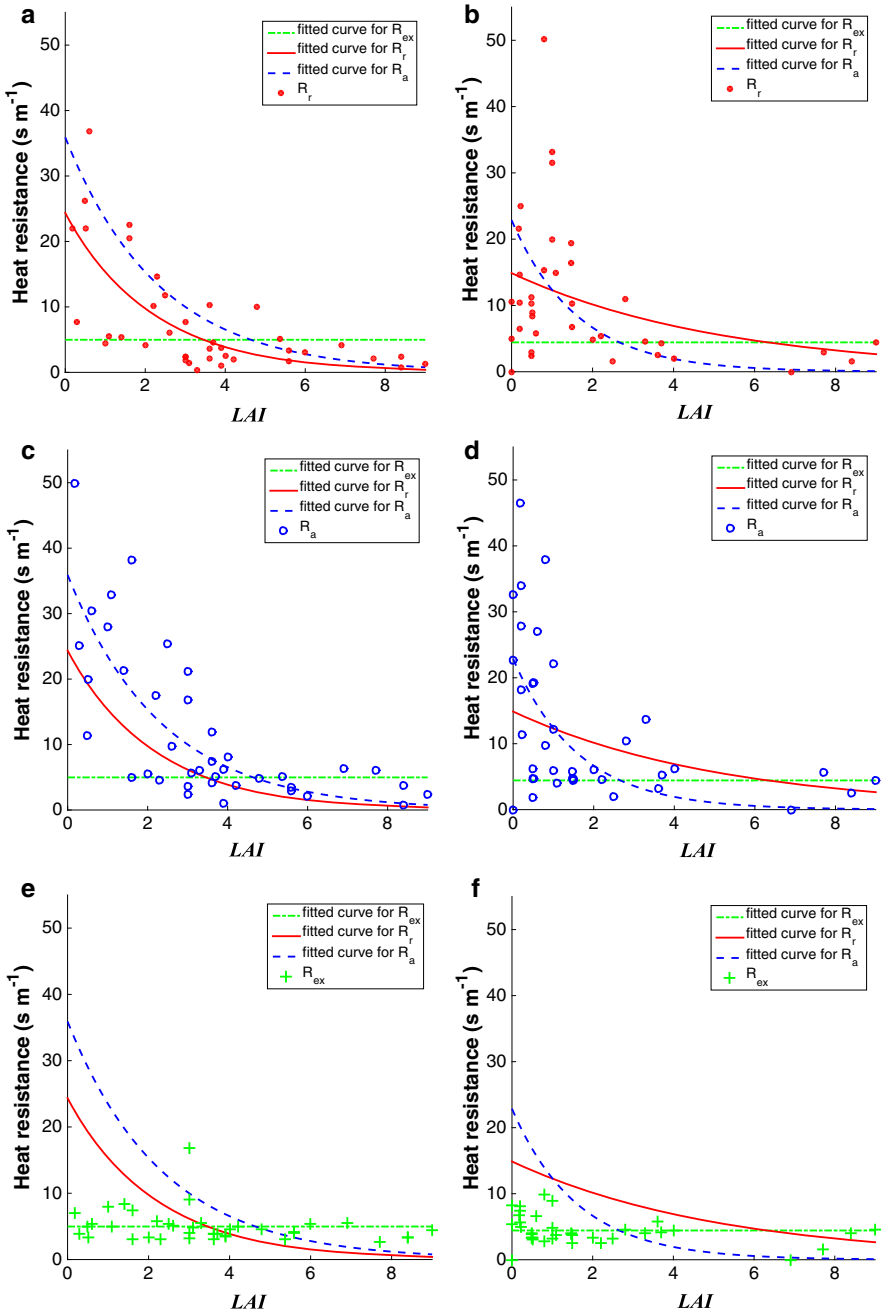
Fig. 2 Relationship between the three heat resistances and *LAI* under unstable conditions. **a**, **c**, **e** summer; **b**, **d**, **f** winter. Each date point represents a site seasonal mean value. *Lines* are regression fits to the data: R_r, $y = 24.4\exp(-0.5x)$ (summer), $y = 14.9\exp(-0.2x)$ (winter); R_a, $y = 35.9\exp(-0.4x)$ (summer), $y = 22.9\exp(-0.6x)$ (winter); R_{ex}, $y = 5\,\mathrm{s\,m^{-1}}$ (summer), $y = 4\,\mathrm{s\,m^{-1}}$ (winter). For each stability class, the same fit *lines* are shown in the three *panels*. The first coefficient in the exponential functions has the dimensions of $\mathrm{s\,m^{-1}}$

Table 2 Statistics of the three component resistances among all the sites

Veg. type	Season	Stability	LAI	R_r	R_a	R_{ex}
				Mean ± s.e. (s m^{-1})	Mean ± s.e. (s m^{-1})	Mean ± s.e. (s m^{-1})
Deciduous	Summer	Unstable	4.1	4.3 ± 09	7.6 ± 2.1	5.5 ± 1.2
		Stable		46.8 ± 6.9	54.8 ± 7.6	9.5 ± 0.8
	Winter	Unstable	1.5	8.6 ± 1.8	9.3 ± 3.1	4.1 ± 0.5
		Stable		78.3 ± 9.7	28.1 ± 5.3	6.1 ± 0.6
Conifer	Summer	Unstable	4.0	6.4 ± 1.8	8.2 ± 2.5	4.1 ± 0.2
		Stable		51.8 ± 6.3	66.5 ± 6.0	10.6 ± 0.8
	Winter	Unstable	3.4	9.3 ± 2.0	9.2 ± 2.4	3.8 ± 0.4
		Stable		59.8 ± 9.4	44.0 ± 5.9	7.3 ± 0.8
Grass/shrub	Summer	Unstable	1.2	17.5 ± 5.2	27.7 ± 3.7	6.0 ± 0.5
		Stable		31.8 ± 1.6	62.6 ± 8.3	8.8 ± 0.9
	Winter	Unstable	0.3	16.9 ± 7.5	20.6 ± 5.7	5.4 ± 1.4
		Stable		56.0 ± 17.8	44.2 ± 11.9	7.1 ± 1.8
Crop	Summer	Unstable	3.6	8.5 ± 0.5	19.8 ± 1.7	6.3 ± 1.0
		Stable		98.3 ± 8.5	58.3 ± 5.2	8.8 ± 0.8
	Winter	Unstable	0.4	17.5 ± 4.8	22.7 ± 1.8	7.4 ± 0.6
		Stable		98.1 ± 16.2	50.4 ± 2.4	8.9 ± 0.1

4.1, the average R_r is 37 s m^{-1}. It appears that the seasonal LAI change alters the radiative property of the surface that is related to the radiometric resistance, but the exact nature of this property is not known.

For the coniferous forest site group, because the difference in the average LAI between the summer and the winter season is very small, the seasonal difference in R_r is greatly reduced (Table 2). The LAI difference is not zero between summer and winter for this group because of the presence of deciduous understorey vegetation and deciduous trees at some of the coniferous sites.

The R_r in the cropland group behaves similarly to in the deciduous forest group, showing much larger values in the winter than in the summer. The grassland/shrubland group, unlike the deciduous forest and cropland groups, shows little difference in R_r between summer and winter under unstable conditions. This is because, for grassland, unless it is burned, dead vegetation is still standing at the surface, whereas for cropland there is a drastic difference in the vegetation stands between the growing and the nongrowing season.

3.3 Daily Variations

To further investigate the influence of LAI, we analyzed the daily variations of the radiometric resistance for a cropland site with continuous maize plantation, Mead irrigated site in Nebraska (site ID US-Ne1, Table 1). The growing season is from the beginning of June to the end of September. This irrigated cropland site has a more variable LAI over the year compared with a typical forest or grassland site.

The patterns seen in the seasonal mean values across multiple sites still hold for the daily values at this specific site (Fig. 3). The multiyear data show obvious annual cycles of R_r, in the range from −20 to 330 s m^{-1} (Fig. 3a, b). As expected, the trend of the R_r cycle is just opposite
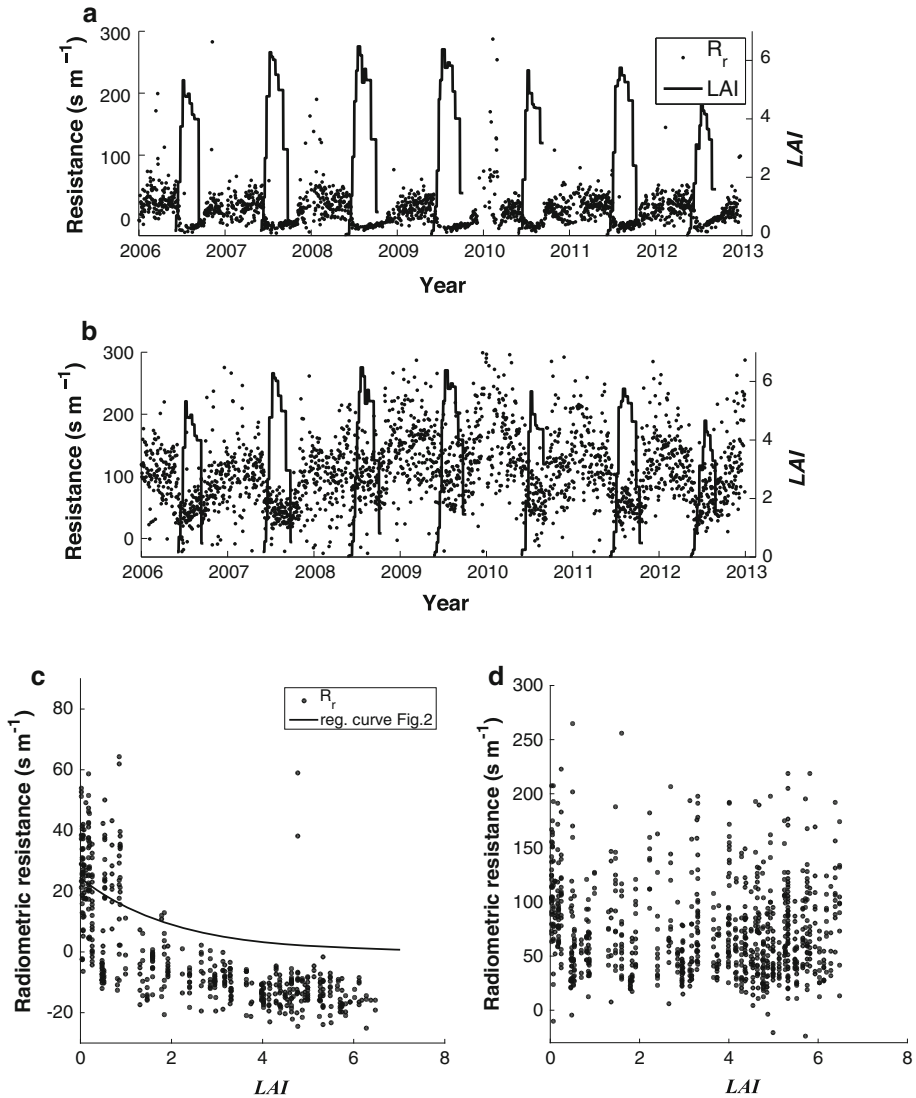
Fig. 3 Correlation between the daily radiometric resistance and *LAI* at the Mead irrigated site in Nebraska (site ID US-Ne1): **a** time series under unstable conditions (daytime), **b** time series under stable conditions (nighttime), **c** scatter plot under unstable conditions, **d** scatter plot under stable conditions. Each *dot* in **a** and **c** represents an average value of all hourly data of a given day for which $H > 5\,\mathrm{W\,m^{-2}}$; each *dot* in **b** and **d** represents an average value of all hourly data of a given day for which $H < -5\,\mathrm{W\,m^{-2}}$. The *regression line* is reproduced from Fig. 2 and represented by $y = 24.4\exp(-0.5x)$

to the phase of the crop *LAI* cycle. The seasonal pattern is stronger under unstable conditions than under stable conditions, as is illustrated by the much clearer variations between the growing and nongrowing season in Fig. 3a than Fig. 3b. This is consistent with what was observed for the seasonal mean values at multiple sites (Fig. 2).

Under unstable conditions, the day-to-day variation is smaller than the contrast between the growing and the nongrowing season; For example, in the year 2008–2009, the standard devi-

ation of the daily R_r is $6\,\mathrm{s\,m^{-1}}$ for the growing season (10 June 2008–7 October 2008) under unstable conditions. For comparison, the mean difference between the two seasons is $57\,\mathrm{s\,m^{-1}}$ under unstable conditions. The day-to-day fluctuations were caused by variations in the site microclimatic conditions. Since the goal of this research is to better isolate the influences on the heat resistances of different land-surface types, these variations represent unwanted "random noise." By using seasonal mean values, these fluctuations were filtered out.

Under unstable conditions, the daily mean R_r shows a pattern (Fig. 3c) similar to the seasonal mean across the sites (Fig. 2a), that is, an exponential decay of R_r with increasing *LAI*. Although the general pattern is similar to that seen for the seasonal mean values at multiple sites, the regression curve is different. The difference is caused by the negative radiometric resistance values that occurred persistently in the growing season (Fig. 3a, c). The two outliers in Fig. 3c probably result from measurement errors. These two dates are 12 July 2006 and 6 July 2007. For both days, only one valid hourly observation is available for calculating the resistance because, for all other hourly measurements, the sign of $\theta_s - \theta_a$ is opposite to the sign of H. Under stable conditions, this exponential decay relationship is no longer detectable (Fig. 3d). Instead, R_r appears independent of *LAI*. The mean R_r value is $105\,\mathrm{s\,m^{-1}}$.

Negative resistances are physically unacceptable. Two possible causes of the negative R_r values at this site are contamination of the air temperature measurement by solar heating, and inaccurate z_0 parametrization. If θ_a is measured correctly, $\theta_s - \theta_a$ should be approximately zero when H is $0\,\mathrm{W\,m^{-2}}$. However, the scatter plot of half-hourly H versus $\theta_s - \theta_a$ for this site (Supplementary Fig. S1) reveals that, at $H = 0$, the temperature difference $\theta_s - \theta_a$ has an offset of $1 - 1.5\,\mathrm{K}$, indicating a potential bias in the temperature measurement. One source of bias error is related to sunlight heating of the air temperature sensor. Another possible source of error is the temperature calibration. The manufacturer's original calibration, not necessarily still accurate, was used at this site. Indeed, reducing θ_a by $1\,\mathrm{K}$ would eliminate most of the negative R_r values and bring the result into better agreement with the site mean regression curve (Fig. 4a).

Raupach (1994) showed that z_0 should respond dynamically to both canopy height and *LAI*. We thus parametrized z_0 as a function of canopy height and *LAI* according to Raupach (1994)
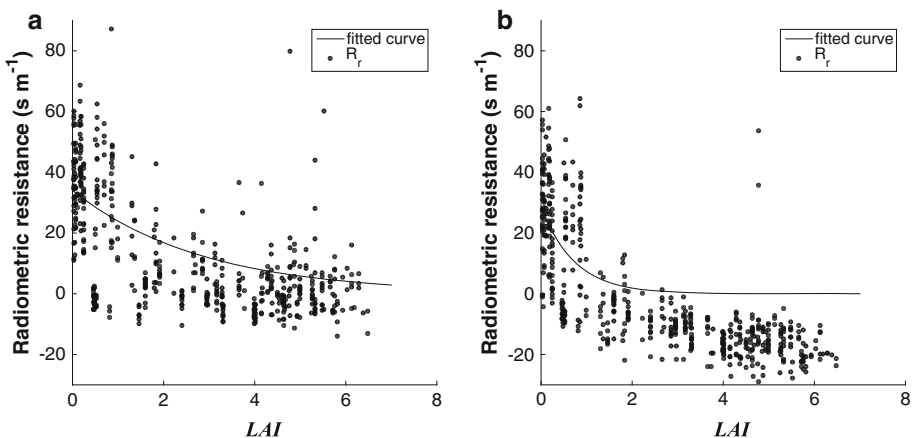

Fig. 4 Sensitivity analyses to air temperature bias and z_0 parameterization under unstable conditions for the Nebraska site: **a** θ_a reduced by $1\,\mathrm{K}$; **b** original θ_a, $z_0 = \exp(-2.3)hLAI^{-0.15}$

to determine whether the negative R_r values may be caused by inaccurate z_0 parametrization. Comparison of this new result (Fig. 4b) with the calculation using fixed z_0 (Fig. 3c) shows that this dynamic parametrization has little effect on the negative R_r values.

3.4 Relations of Heat Resistances to Thermal Stability

All the three component resistances are generally larger under stable conditions than under unstable conditions (Figs. 2, 5). Table 2 presents the average resistances across all sites under unstable and stable conditions and in the summer and the winter. Our results confirm that stable stratification tends to suppress heat transfer. This thermal stability effect has been known for a long time for aerodynamic resistance (Garratt 1994). According to our results, thermal stability has similar impacts on the radiometric resistance. Table 2 shows that R_r is larger under stable conditions than under unstable conditions in each vegetation type group, and in both winter and summer. On average, summer R_r is 68 s m^{-1} larger, and winter R_r is 64 s m^{-1} larger under stable conditions than under unstable conditions for all the selected sites. It appears that stable stratification enlarges the distance between z_h and z_r.

Under stable conditions, there is no apparent correlation between *LAI* and the aerodynamic resistance (summer: $R^2 < 0.05$; winter: $R^2 < 0.01$), a pattern that is different from that observed in unstable conditions (summer: $R^2 > 0.58$; winter: $R^2 > 0.30$). Under stable conditions, R_r shows no apparent correlation with *LAI* either (Fig. 5a, b), and is on average 40 s m^{-1} larger than R_{ex} in the summer and 65 s m^{-1} larger in the winter for all the sites examined in this study. The R_r values are compatible with the R_r estimates by Stewart et al. (1994) (22–73 s m^{-1}). Under unstable conditions, the average R_r is 8 s m^{-1}, only slightly larger than R_{ex} (5 s m^{-1}).

3.5 Impacts of Radiometric Resistance on Sensible Heat Flux Calculation

To assess the impacts of radiometric resistance on the sensible heat flux calculation, we first calculated the sensible heat flux using the complete Eq. 6 and again by omitting R_r from Eq. 6, as

$$\hat{H} = \rho C_p \frac{\theta_s - \theta_a}{R_a + R_{ex}}. \tag{11}$$

In Eq. 6, R_r was obtained from the regression fit functions of *LAI* shown in Fig. 2a (summer) and 2b (winter) for unstable conditions and constant values of 50 s m^{-1} (summer) and 66 s m^{-1} (winter) for stable conditions, and other terms were provided by field measurements. The prediction error e and the relative prediction error e_r are defined as

$$e = \hat{H} - H, \tag{12}$$

$$e_r = \frac{e}{H}, \tag{13}$$

where H is the observed sensible heat flux.

Unsurprisingly, omission of R_r results in an overestimation of the magnitude of the sensible heat flux. Under stable conditions, the prediction error has no apparent correlation with the site *LAI*. Under unstable conditions, the prediction error is much larger for low-*LAI* sites than for high-*LAI* sites (Fig. 6a). The relative error decreases as the *LAI* increases under unstable conditions and has no apparent correlation with *LAI* under stable conditions (Fig. 6b). These results show that the largest overestimation due to omission of R_r occurs at low-*LAI* surfaces and under unstable conditions.
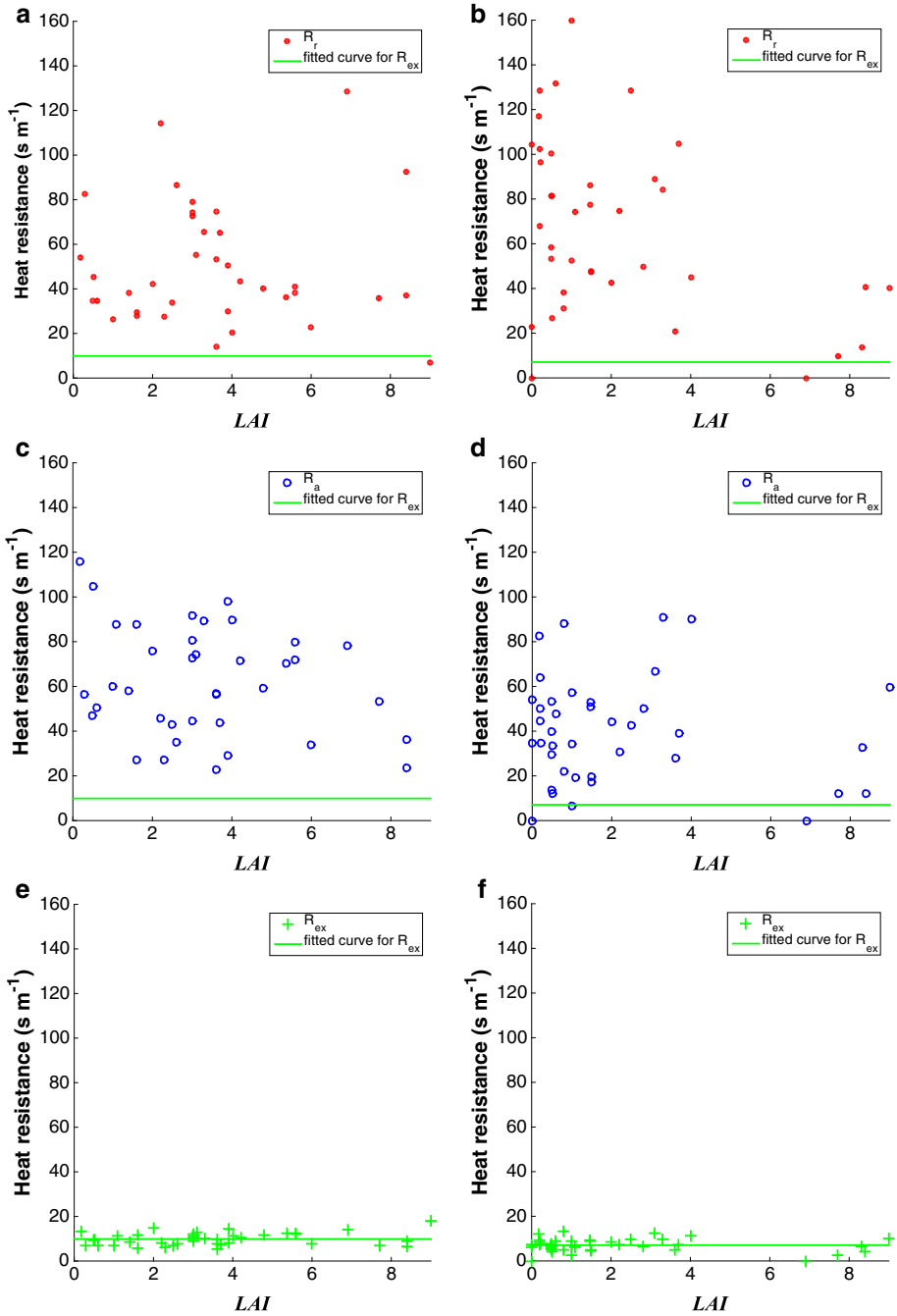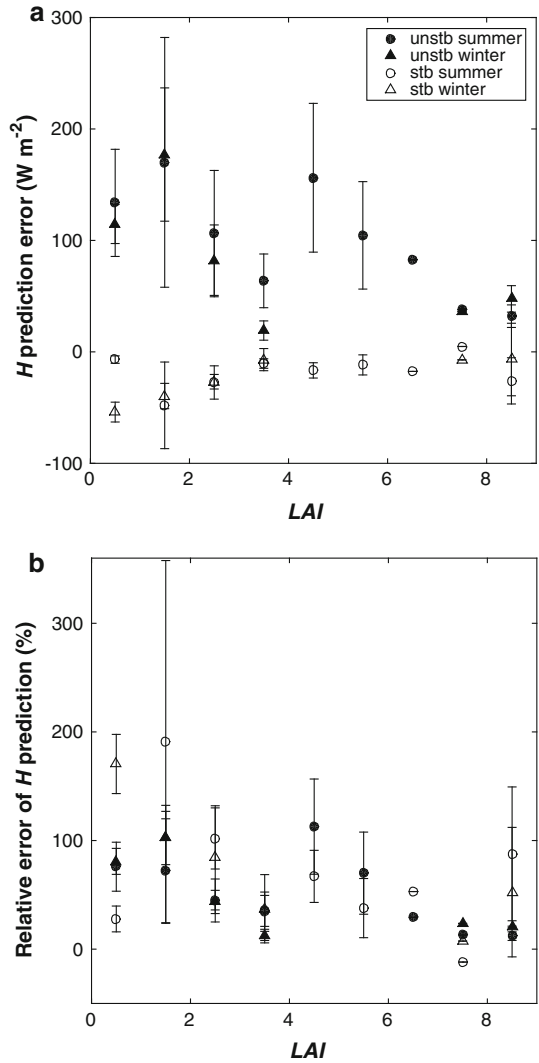
Fig. 5 Same as Fig. 2 but for stable conditions

Fig. 6 Bin average prediction errors of sensible heat flux caused by omission of the radiometric resistance: *top panel* prediction error; *bottom panel* relative prediction error. *Error bars* denote 1 s.e.

Similar results are found in the literature. In an experiment in a soybean field, Lee et al. (2009) reported larger prediction errors (up to $308\,\mathrm{W\,m^{-2}}$) for the early part of the growing season ($LAI < 2$) than in the middle part of the growing season (error $< 15\,\mathrm{W\,m^{-2}}$) when the *LAI* is large ($LAI = 7.6$). Chen and Zhang (2009) compared the observed heat transfer coefficient C_h from Ameriflux sites and the modeled C_h according to the Noah LSM, and reported large model overestimations for short vegetation. In the LSM, the two roughness values (z_h and z_0) are used to calculate C_h without the radiometric resistance. Because C_h is inversely proportional to the heat resistance, overestimation of C_h indicates overestimation of H in the model. Zheng et al. (2012) reported a large cold bias in the daytime θ_s (up to $-15\,^{\circ}\mathrm{C}$) produced by NCEP Global Forecast Modelling System operational modeling over the arid western continental USA, where *LAI* is very small. In the Global Forecast System, θ_s is solved from the energy balance equation with H parametrized according to Eq. 1 and with the

approximation $\theta_s = \theta_{0h}$. In other words, the default NCEP calculation omits the radiometric resistance. By adding the radiometric resistance as a function of vegetation fraction to the bulk formula, Zheng et al. (2012) significantly reduced the cold bias in θ_s (average bias $< -4\,°C$). Similarly to our findings, they found that their R_r formulation had a minimal effect on the nighttime θ_s under stable conditions. Recently, Zhang et al. (2014) also reported a large high bias ($120\,W\,m^{-2}$) in H calculated with a *LAI*-independent parametrization of z_r for a desert steppe site, and using a time-varying roughness length for heat as a function of vegetation growth reduced the root-mean-squared error of H by more than half.

4 Conclusions

A simple resistance model was used to decompose the total resistance to heat transfer from the surface to a reference height into three additive components: radiometric resistance, excess resistance, and aerodynamic resistance. On average, all resistances are greater under stable conditions than under unstable conditions. The excess resistance shows no apparent relation to *LAI* changes under both unstable and stable conditions. The aerodynamic and radiometric resistances decrease exponentially as *LAI* increases under unstable conditions. Under stable conditions, the aerodynamic and radiometric resistances show no apparent relation to *LAI* changes and are much larger than the excess resistance. Daily data from a cropland site show that the correlations with *LAI* seen in the seasonal mean values across multiple sites still hold for the daily values. High bias in the sensible heat flux calculation occurs if the radiometric resistance is omitted from the bulk transfer formulation, with the overestimation larger at lower *LAI*. This problem is especially severe at low *LAI* (< 2) and under unstable conditions.

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